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2 **Supporting information for**
3 Secondary Organic Aerosol from OH Oxidation of Acyclic Terpenes is More
4 Viscous and Less Volatile Than That of Their Cyclic Analogs

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14 **Table S1.** The experimental conditions and particle size selected for evaporation
15 kinetics.

Experiment date	VOC type	Size selected for evaporation kinetics (d _m , nm)	Calculated density (g/cm ³)	Targeted VOC concentration (ppb)	GC-estimated VOC concentration (ppb)
08/23/22	α-pinene	165	1.24 ± 0.01	285	N/A
08/25/22	β-ocimene	220	1.35 ± 0.01	283	N/A
09/01/22	β-ocimene	150		50	N/A
09/06/23	β-caryophyllene	165	1.17 ± 0.01	53.4	N/A
09/07/23	β-caryophyllene	130	1.16 ± 0.01	32.0	15.0
09/08/23	β-farnesene	140	1.20 ± 0.01	48.2	N/A
09/11/23	β-farnesene	135	1.19 ± 0.01	40	11.7

16 **Nano-DESI-HRMS sample and MFAssignR analysis.**

17 Due to the limited collected SOA mass, we utilized nano-DESI-HRMS for offline
18 compositional analysis. A general description of the nano-DESI instrument can be found
19 in Roach *et al.*, 2010.¹ The nano-DESI interface capillary (Polymicro Technology L.L.C.,
20 Phoenix, USA) was coupled to the inlet of the mass spectrometer, with two capillaries
21 aligned at a 90° angle. The capillary features an outer diameter of 150 μm and an inner
22 diameter of 50 μm. The secondary capillary, approximately 1.5 cm in length, was

23 positioned roughly 1 mm away from the MS inlet. A solvent mixture of 7:3 acetonitrile:
24 water (Optima LC-MS grade, Fisher Chemical, USA) was delivered at a flow rate of 500
25 nL/min through the primary capillary by a syringe pump. This arrangement facilitated the
26 formation of a liquid interface between the capillaries on the sample surface, allowing
27 for the continuous transfer of extracted analytes to the orbitrap spectrometer. A high
28 voltage of -3.5 kV was applied to the syringe needle in negative ion mode. The capillary
29 temperature was maintained at 275°C. The maximum ion injection time and the
30 Automatic Gain Control (AGC) target were set to 500 ms and 5×10^5 , respectively. The
31 mass spectrometer was configured to acquire mass spectra from m/z 100-1300 at a
32 mass resolution of 100,000 (at m/z 400). Via a XYZ Zaber sample stage and custom
33 LabVIEW software, the sample was scanned along the XY plane under the nano-DESI
34 liquid junction at 75 $\mu\text{m/s}$, ensuring consistent and fresh sample desorption for
35 consecutively collected HRMS scans (100 scans).

36 MFAssignR was employed for assigning formulas to detected m/z values;²
37 comprehensive details about the software are provided in previous publications.^{3,4}
38 Noise and isotope features were filtered out from the mass list, with the final
39 assignments to formulae $\text{C}_x\text{H}_y\text{O}_z$ being made with an accuracy of 4 ppm. The
40 processing of both blank and sample data followed a uniform protocol. Subsequently,
41 signals present in both the blank and sample were removed if the signal in the sample
42 was not at least 10 times greater than that in the blank. We analyzed both Teflon filter
43 blanks and solvent blanks using the same nano-DESI protocol as for the samples.
44 Peaks present in the blanks were removed unless the corresponding signal intensity in
45 the sample was at least an order of magnitude ($\geq 10\times$) higher than in the blank. Only
46 these blank-screened peaks were retained for subsequent data analysis and
47 interpretation.

48 Chemodiversity was estimated using Shannon's index (H):⁵

49
$$H = - \sum_{i=1}^S p_i \times \ln p_i$$

50 and in this formula, p_i represents the normalized intensity of individual formula product,
51 and S is the species richness.

52 **7-bin Evaporation Kinetics Fits and Modeling.**

53 The methods used to fit and model evaporation kinetics have been described in detail
54 elsewhere and are summarized here.⁶⁻⁸ Measured evaporation data were fit with a bi-
55 exponential decay function of the form $\text{VFR} = \text{Ae}^{-at} + \text{Be}^{-bt}$, where t is the evaporation
56 time. In brief, time-dependent evaporation of multicomponent single particles was
57 simulated using a seven-bin VBS, assuming that activated charcoal removes gas-phase

58 organics as they evaporate. Particles were treated as uniform and non-interacting. The
 59 initial kinetic equations simulate single-particle evaporation given an assumed VBS, and
 60 the VBS mass fractions are iteratively adjusted until the modeled kinetics match the
 61 measurements. Volatility bins span effective saturation concentrations from 1×10^{-4} to
 62 1×10^2 . The lower boundary for the lowest-volatility bin was defined at this value because
 63 the compounds with this and lower effective saturation concentrations remain in the
 64 particle phase at room temperature within the timescale of the experiments (~ 24 hours).
 65 This means we are unable to resolve compounds in different bins at volatilities lower
 66 than this even though other papers have defined ELVOC as $c^* < 3 \times 10^{-5}$.

67

68 **Viscosity and Volatility Estimation.**

69 Glass transition temperature (T_g) is the point at which the state of SOA changes from
 70 viscous semisolid to a glassy solid state, and T_g of SOA can be predicted using the
 71 molecular formulae from the mass spectrometry:⁹

$$72 \quad T_g = (n_C^0 + \ln(n_C))b_C + \ln(n_H)b_H + \ln(n_C)\ln(n_H)b_{CH} + \ln(n_O)b_O + \ln(n_C)\ln(n_O)b_{CO},$$

73 where n_C^0 stands for reference carbon number, while b_C , b_H , b_{CH} , b_O , and b_{CO}
 74 represent the contribution of carbon, hydrogen, carbon-hydrogen interaction, oxygen,
 75 and carbo-oxygen interaction to T_g , and the values are listed in Table S2.

76 Chemical compound classes and parameters for estimating glass transition
 77 temperature, as described by DeRieux et al.⁹

Classes	n_C^0	b_C	b_H	b_{CH}	b_O	b_{CO}
CH	1.96	61.99	-113.33	28.74		
CHO	12.13	10.95	-41.82	21.61	118.96	-24.38

78

79 Under dry conditions, the T_g values for a mixture of SOA were derived using a linear
 80 approach through the Gordon-Taylor method, which assumes Gordon-Taylor constant
 81 (k_{GT}) equal to 1, as expressed by the formula below.^{10,11}

$$82 \quad T_g = \sum_i w_i T_{g_i}$$

83

84 Here, w_i indicates the mass fraction of the i^{th} component, derived from mass spectra
 85 peak abundances. In humid conditions, the presence of water alter the T_g , which can be
 86 calculated for a water-organic mix using Gordon–Taylor equation:¹²

87
$$T_{g(w_{org})} = \frac{(1 - w_{org}) T_{g,w} + \frac{1}{k_{GT}} w_{org} T_{g,org}}{(1 - w_{org}) + \frac{1}{k_{GT}} w_{org}}$$

88

89 where w_{org} is the mass fraction of organic components, $T_{g,w}$ is the glass transition
 90 temperature of water (136 K), and k_{GT} is the Gordon-Taylor constant, set here to
 91 2.5.^{12,13}

92

93 The effective hygroscopicity parameter (κ) was assumed to be 0.10 for monoterpene-
 94 derived SOA and 0.05 for sesquiterpene-derived SOA due to the previously observed
 95 hygroscopicity differences and estimated values,¹⁴⁻¹⁹ consistent with values reported in
 96 previous studies. These values were used in the computation of the mass
 97 concentrations of SOA (m_{SOA}) and water (m_{H_2O}):

98
$$m_{H_2O} = \frac{\kappa \rho_w m_{SOA}}{\rho_{SOA} \left(\frac{1}{a_w} - 1 \right)}$$

99 Here, ρ_w and ρ_{SOA} denote the densities of water and SOA, measured at 1 g cm⁻³ and
 100 1.4 g cm⁻³ respectively, whereas a_w symbolizes water activity, which is computed as
 101 $a_w = 100/RH$.

102 Furthermore, the dependence of viscosity on temperature was determined through the
 103 Vogel-Tammann-Fulcher (VTF) equation:

104
$$\log(\eta) = -5 + 0.434 \frac{T_0 D_f}{T - T_0}$$

105 In this equation, T_0 stands for the Vogel temperature, which can be calculated using T_0
 106 $= \frac{39.17 T_g}{D_f + 39.17}$, D_f is the fragility parameter, quantifies the deviation from Arrhenius
 107 behavior and was postulated to be 10.⁹

108 To calculate the volatility distribution of SOA, the parameterization approach by Li et al.
 109 was employed.²⁰ This method allows for the estimation of the saturation mass
 110 concentrations (C_0) of pure compounds using the formula:

111
$$\log_{10} C_0 = (n_c^0 - n_c) b_c - n_o b_o - 2 \frac{n_c n_o}{n_c + n_o} b_{CO}$$

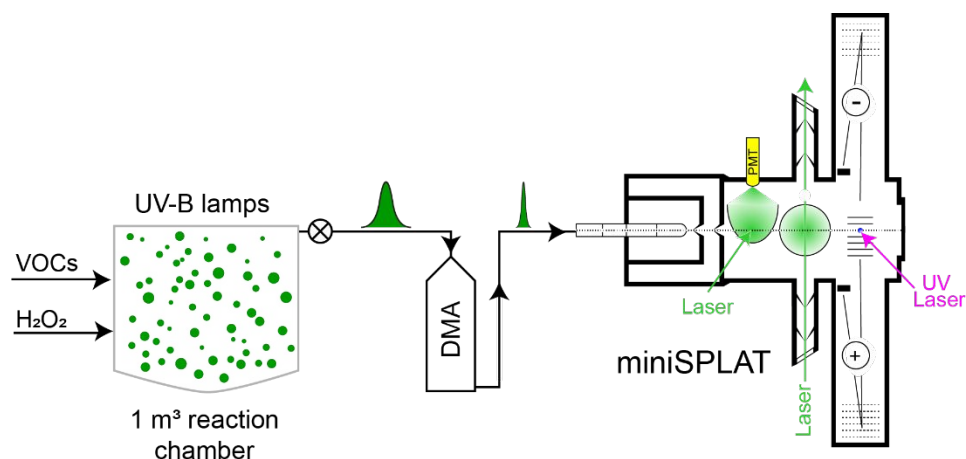
112 Here, n_{CO} represents the reference carbon number, while n_c and n_o indicate the counts
 113 of carbon and oxygen atoms, respectively. The coefficients b_c , b_o , and b_{CO} are provided
 114 in Table S3. Although this formulaic approach is suited for estimating volatility in

115 isomeric mixtures, it's important to note that the absence of detailed structural
116 information introduces a degree of uncertainty to these estimations.

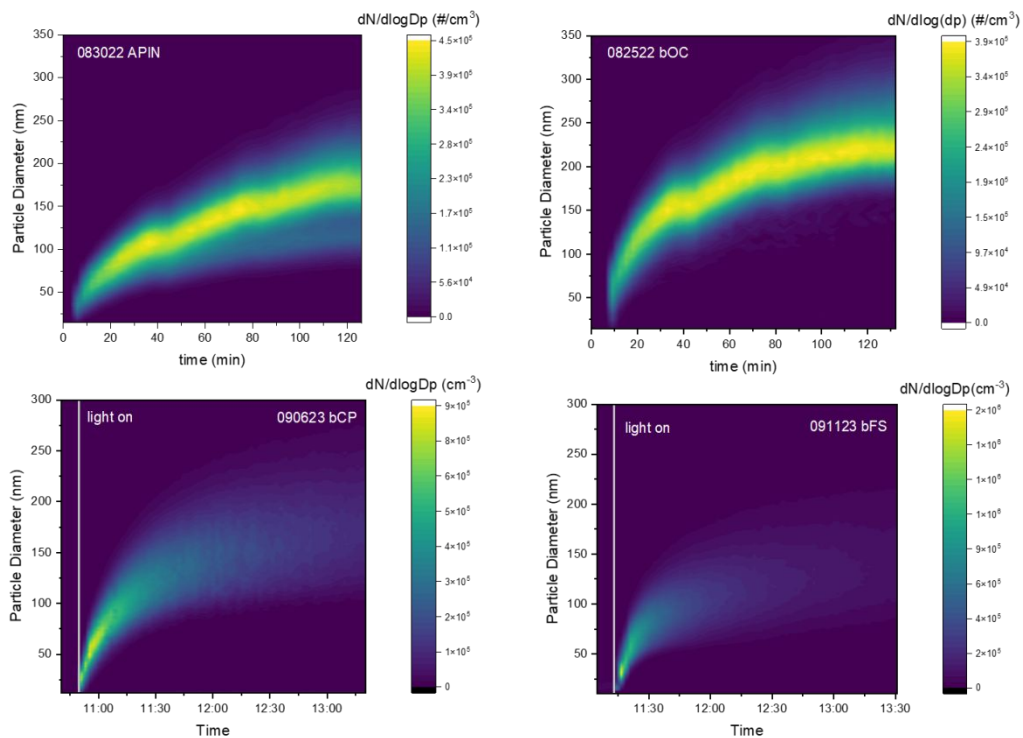
117 **Table S3.** Classes of chemical compounds and saturation mass concentration
118 parameterization described by Li et al..²⁰

Classes	n_c^0	b_c	b_o	b_{co}
CH	23.8	0.48861		
CHO	22.66	0.4481	1.656	-0.7790

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125 **Figure S1.** Schematic of the experimental setup for miniSPLAT measurements of
126 density and shape of size-selected particles generated in the reaction chamber by
127 photooxidation.



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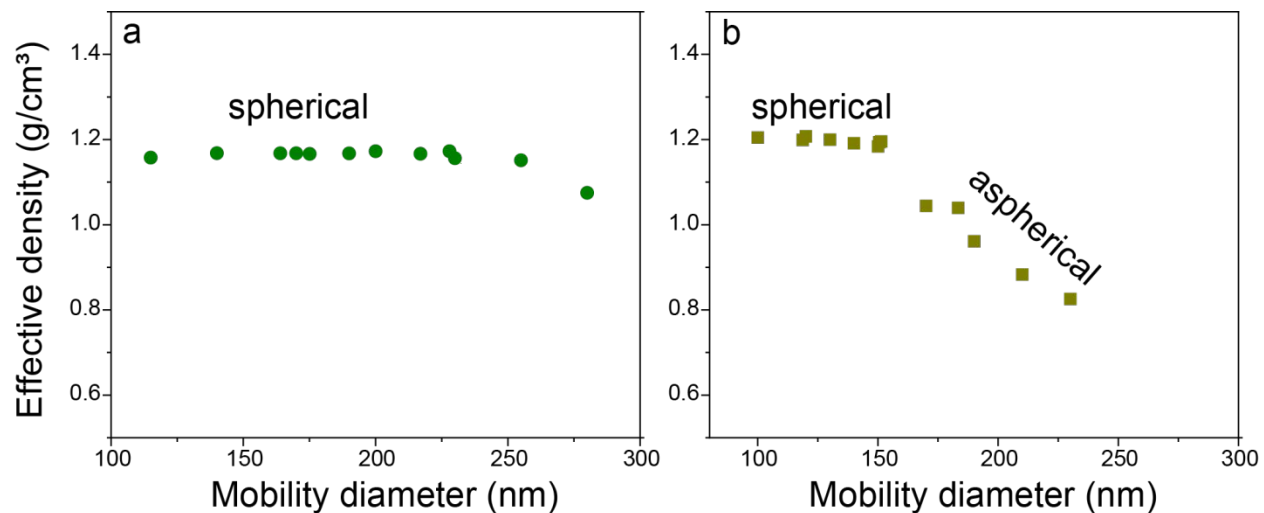
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130 **Figure S2** The observed evolution of particle mobility diameter during SOA growth, with
 131 color representing the $dN/d\log D_p$. The small gaps observed for monoterpene SOA
 132 resulted from VOC cartridges sampling, with the UV-lights being temporarily off.

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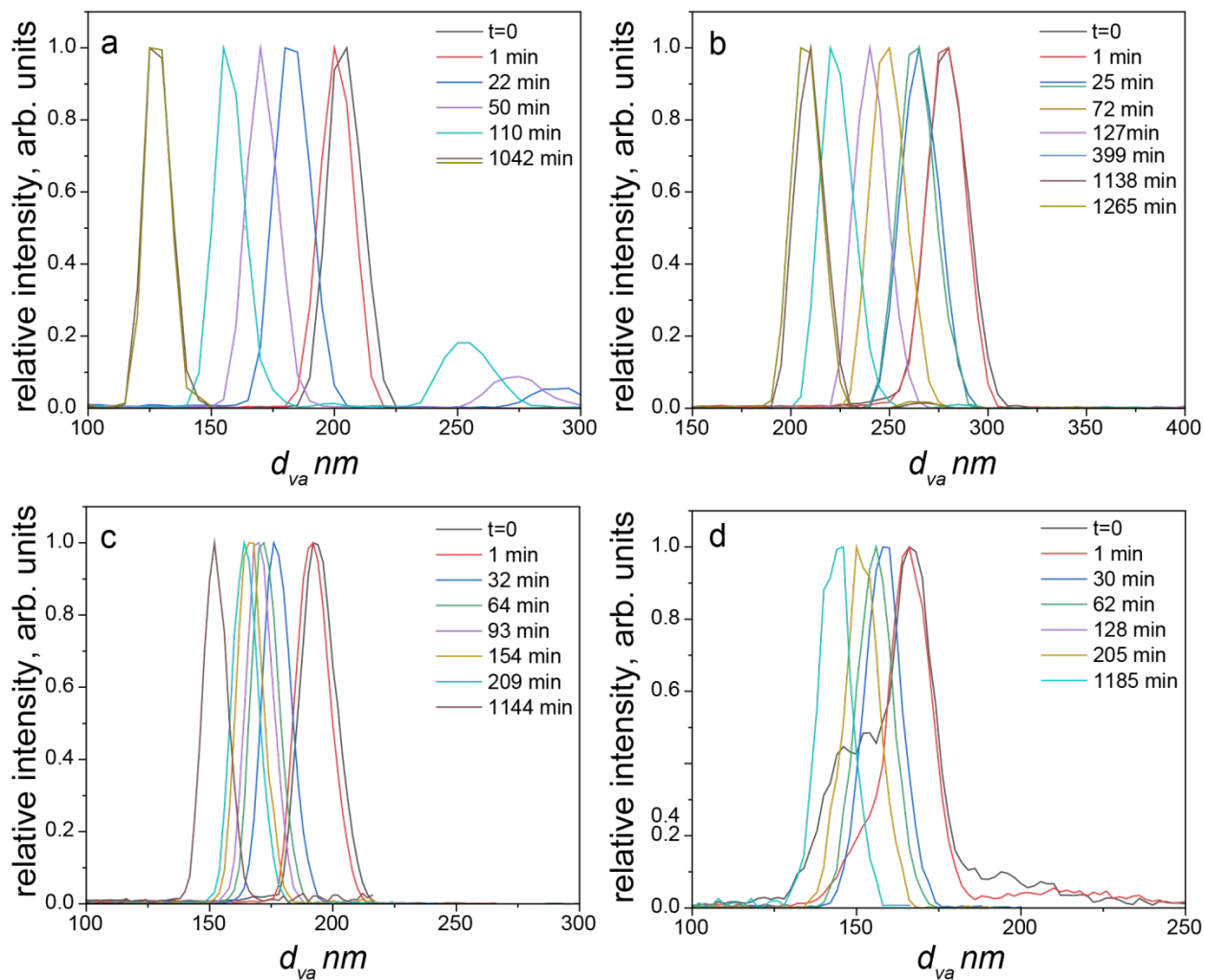
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137 **Figure S3** The relationship between density/effective density and mobility diameter (d_m)
 138 for (a) β -caryophyllene, and (b) β -farnesene SOA size-selected particles, indicating the
 139 small particles are spherical, while larger particles become aspherical and have lower
 140 effective densities.

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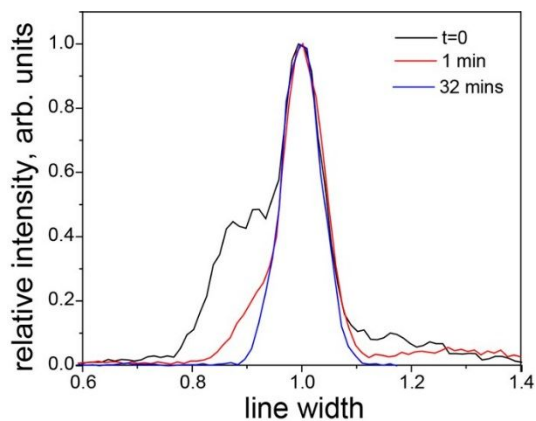


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145 **Figure S4** The d_{va} distribution as a function of evaporation time of (a) α -pinene (doublet
 146 particles observed), (b) β -ocimene, (c) β -caryophyllene, and (d) β -farnesene SOA
 147 during wet evaporation.

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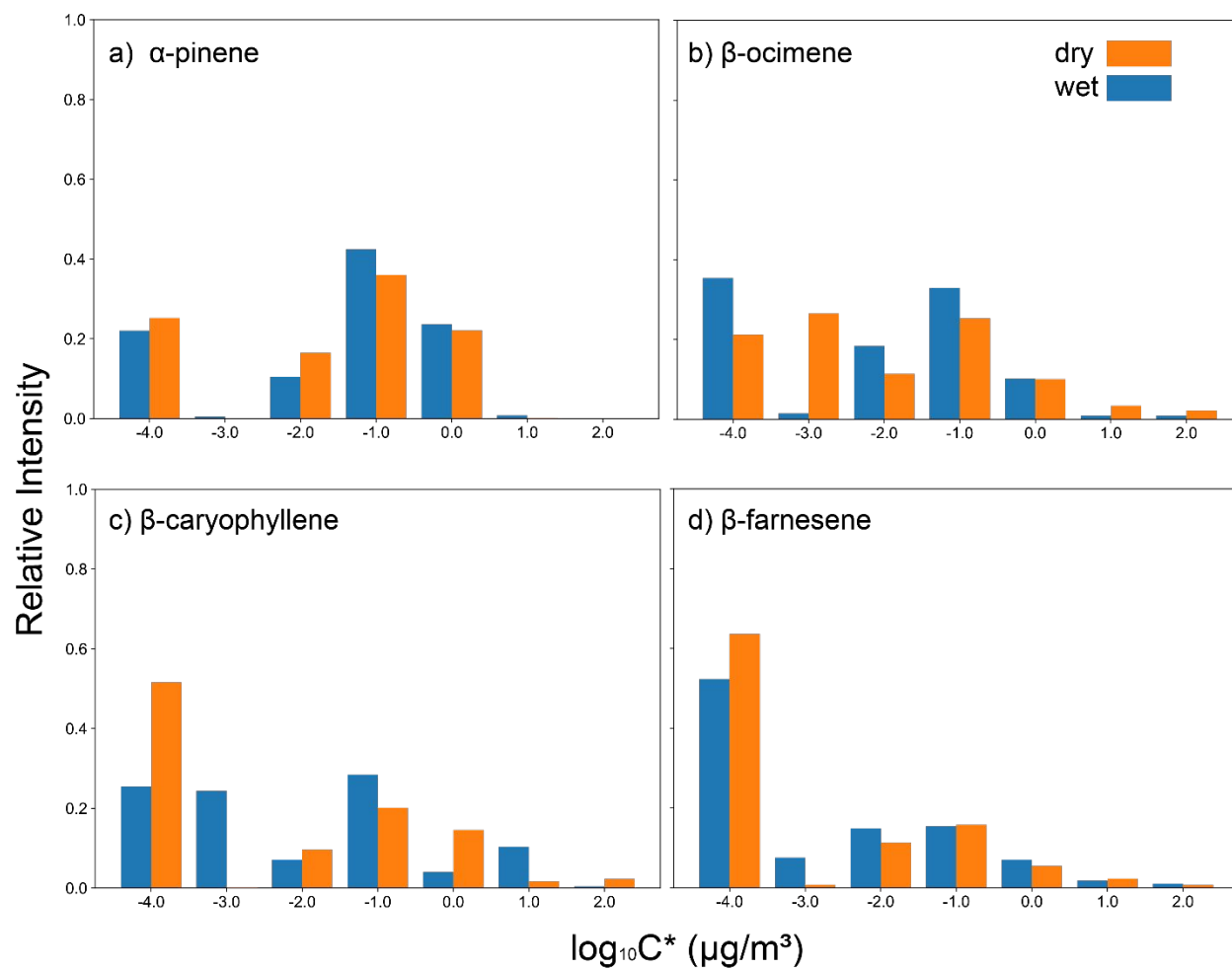
151 **Figure S5** Linewidths of d_{va} distributions for β -farnesene SOA particles measured at 0-,
152 1-, and 32-minutes during evaporation, illustrating changes in particle shape and
153 distribution broadness over time.

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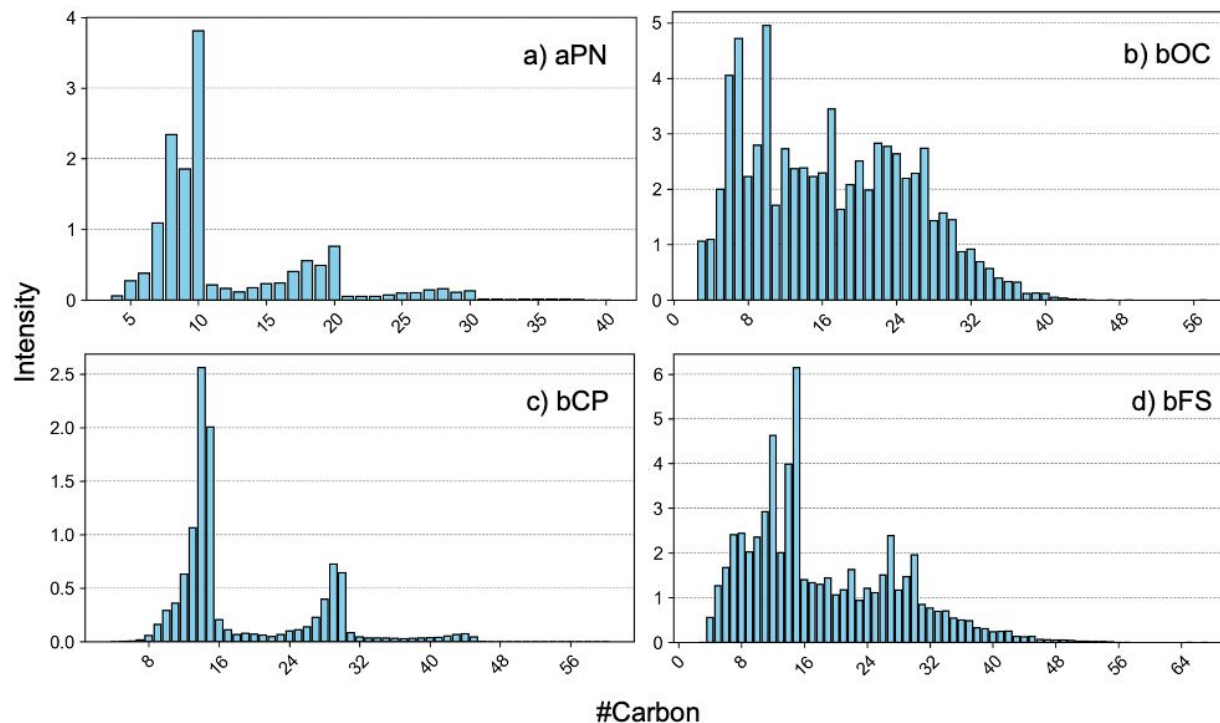
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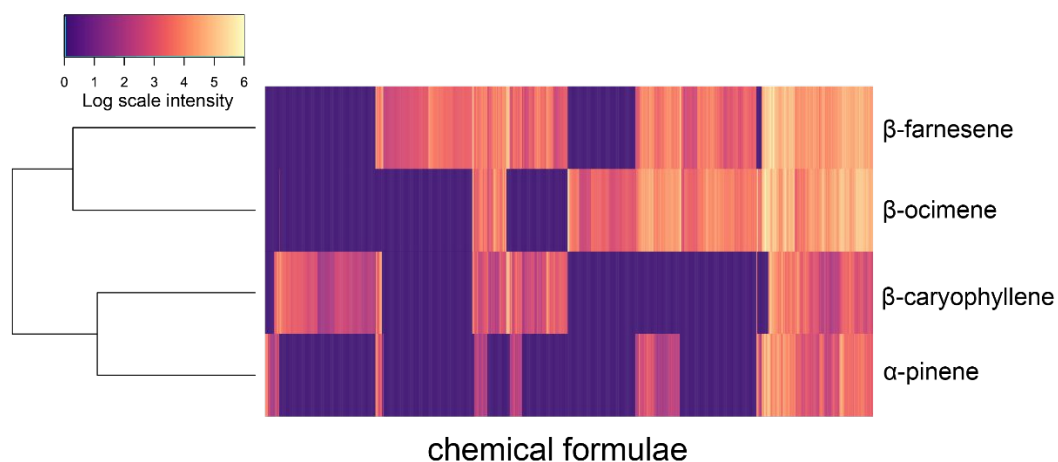
160 **Figure S6** Comparison of wet and dry fraction value across 7 volatility bins on a
 161 logarithmic scale, derived from evaporation kinetics.



162

163 **Figure S7** Total intensity distribution by carbon number for (a) α -pinene(aPN), (b) β -
 164 ocimene(bOC), (c) β -caryophyllene (Bcp), and (d) β -farnesene SOA (bFS). This bar
 165 chart represents the summed intensities for each carbon number, derived from nano-
 166 DESI-HRMS measurements.

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169 **Figure S8** Heatmap comparison of high-resolution mass spectra from four SOA
 170 samples. Each vertical strip corresponds to aligned formula, while the rows represent
 171 individual samples. Each cell corresponds to the signal intensity of a particular formula
 172 in a given sample, with color ranging from low (dark purple) to high (yellow). The
 173 heatmap highlights groups of formulas that co-occur at high abundance in specific

174 samples and help visualize the comparison of the composition of the different SOA
 175 systems. Samples with similar intensities for the same formulas were clustered.
 176 Distances were calculated from the normalized peak-intensity vectors using Euclidean
 177 distance, and clusters were formed using complete-linkage hierarchical clustering.

178

179

180 **Table S4.** Fitted parameters α -values using the 7-bin volatility basis set (VBS) derived
 181 from evaporation kinetics dry experiments for four terpene compounds: α -pinene, β -
 182 ocimene, β -caryophyllene, and β -farnesene. The α -values represent the mass fraction
 183 yields of the oxidation products in each of the c^* bins.

184

C*($\mu\text{g}/\text{m}^3$)	0.0001	0.001	0.01	0.1	1	10	100
α-pinene	2.52E-01	1.57E-04	1.64E-01	3.60E-01	2.21E-01	1.42E-03	1.58E-04
β-ocimene	2.12E-01	2.65E-01	1.14E-01	2.53E-01	1.00E-01	3.36E-02	2.23E-02
β-caryophyllene	5.16E-01	1.95E-03	9.71E-02	2.01E-01	1.45E-01	1.58E-02	2.28E-02
β-farnesene	6.36E-01	7.81E-03	1.13E-01	1.58E-01	5.51E-02	2.29E-02	7.12E-03

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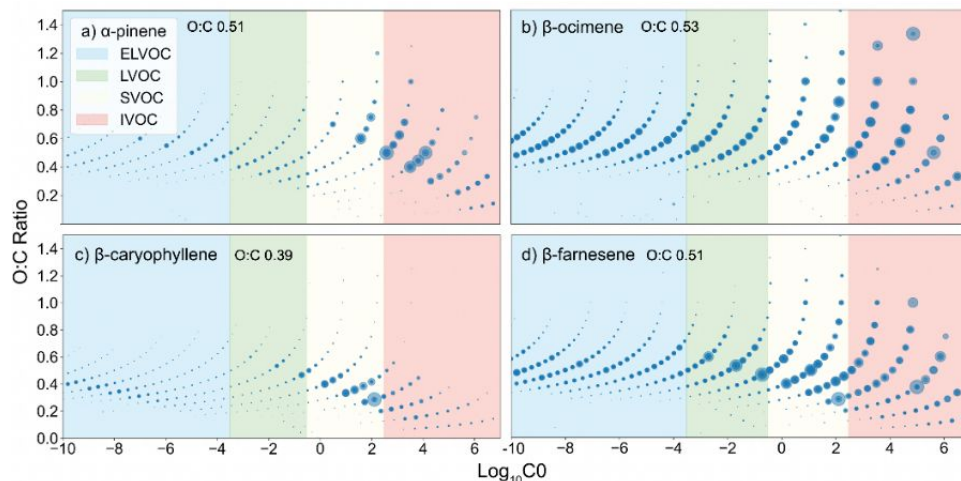
187 **Table S5.** Fitted parameters α -values using the 7-bin volatility basis set (VBS) derived
 188 from evaporation kinetics wet experiments for four terpene compounds: α -pinene, β -
 189 ocimene, β -caryophyllene, and β -farnesene. The α -values represent the mass fraction
 190 yields of the oxidation products in each of the c^* bins.

C*($\mu\text{g}/\text{m}^3$)	0.0001	0.001	0.01	0.1	1	10	100
α-pinene	2.20E-01	4.44E-03	1.05E-01	4.25E-01	2.37E-01	8.05E-03	7.51E-04
β-ocimene	3.54E-01	1.45E-02	1.83E-01	3.29E-01	1.02E-01	8.03E-03	9.00E-03
β-caryophyllene	2.54E-01	2.43E-01	7.09E-02	2.84E-01	3.98E-02	1.04E-01	4.48E-03
β-farnesene	5.23E-01	7.52E-02	1.49E-01	1.54E-01	7.06E-02	1.77E-02	1.02E-02

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195 **Figure S9.** The correlation between the oxygen-to-carbon (O:C) ratio and the volatility,
 196 indicated by the $\log_{10}C_0$. These estimations are based on chemical formulae identified
 197 through nano-DESI-HRMS analysis for SOA originating from OH oxidation of (a) α -
 198 pinene, (b) β -ocimene, (c) β -caryophyllene, and (d) β -farnesene. The bubble size within
 199 the graph reflects the logarithmic scale of the normalized intensity, offering a visual
 200 representation of compound abundance.

201

202 References

- 203 (1) Roach, P. J.; Laskin, J.; Laskin, A. Molecular Characterization of Organic Aerosols
 204 Using Nanospray-Desorption/Electrospray Ionization-Mass Spectrometry. *Anal.*
 205 *Chem.* **2010**, *82* (19), 7979–7986. <https://doi.org/10.1021/ac101449p>.
 206 (2) Schum, S. K.; Brown, L. E.; Mazzoleni, L. R. MFAssignR: Molecular Formula
 207 Assignment Software for Ultrahigh Resolution Mass Spectrometry Analysis of
 208 Environmental Complex Mixtures. *Environ. Res.* **2020**, *191*, 110114.
 209 <https://doi.org/10.1016/J.ENVRES.2020.110114>.
 210 (3) Vandergrift, G. W.; Dexheimer, D. N.; Zhang, D.; Cheng, Z.; Lata, N. N.; Rogers, M.
 211 M.; Shrivastava, M.; Zhang, J.; Gaudet, B. J.; Mei, F.; China, S. Tethered Balloon
 212 System and High-Resolution Mass Spectrometry Reveal Increased Organonitrates
 213 Aloft Compared to the Ground Level. *Environ. Sci. Technol.* **2024**, *58* (23), 10060–
 214 10071. <https://doi.org/10.1021/acs.est.4c02090>.
 215 (4) Vandergrift, G. W.; Shawon, A. S. M.; Dexheimer, D. N.; Zawadowicz, M. A.; Mei, F.;
 216 China, S. Molecular Characterization of Organosulfate-Dominated Aerosols over
 217 Agricultural Fields from the Southern Great Plains by High-Resolution Mass
 218 Spectrometry. *ACS Earth Space Chem.* **2022**, *6* (7), 1733–1741.
 219 <https://doi.org/10.1021/acsearthspacechem.2c00043>.
 220 (5) Shannon, C. E. A Mathematical Theory of Communication. *Bell Syst. Tech. J.* **1948**,
 221 *27* (3), 379–423. <https://doi.org/10.1002/j.1538-7305.1948.tb01338.x>.
 222 (6) Wilson, J.; Imre, D.; Beránek, J.; Shrivastava, M.; Zelenyuk, A. Evaporation Kinetics
 223 of Laboratory-Generated Secondary Organic Aerosols at Elevated Relative

- 224 Humidity. *Environ. Sci. Technol.* **2015**, *49* (1), 243–249.
225 <https://doi.org/10.1021/es505331d>.
- 226 (7) Vaden, T. D.; Imre, D.; Beránek, J.; Shrivastava, M.; Zelenyuk, A. Evaporation
227 Kinetics and Phase of Laboratory and Ambient Secondary Organic Aerosol. *Proc.*
228 *Natl. Acad. Sci.* **2011**, *108* (6), 2190–2195.
229 <https://doi.org/10.1073/pnas.1013391108>.
- 230 (8) Shrivastava, M.; Zelenyuk, A.; Imre, D.; Easter, R.; Beranek, J.; Zaveri, R. A.; Fast,
231 J. Implications of Low Volatility SOA and Gas-Phase Fragmentation Reactions on
232 SOA Loadings and Their Spatial and Temporal Evolution in the Atmosphere. *J.*
233 *Geophys. Res. Atmospheres* **2013**, *118* (8), 3328–3342.
234 <https://doi.org/10.1002/jgrd.50160>.
- 235 (9) DeRieux, W. S. W.; Li, Y.; Lin, P.; Laskin, J.; Laskin, A.; Bertram, A. K.; Nizkorodov,
236 S. A.; Shiraiwa, M. Predicting the Glass Transition Temperature and Viscosity of
237 Secondary Organic Material Using Molecular Composition. *Atmospheric Chem.*
238 *Phys.* **2018**, *18* (9), 6331–6351. <https://doi.org/10.5194/ACP-18-6331-2018>.
- 239 (10) Gordon, M.; Taylor, J. S. Ideal Copolymers and the Second-Order Transitions of
240 Synthetic Rubbers. i. Non-Crystalline Copolymers. *J. Appl. Chem.* **2007**, *2* (9), 493–
241 500. <https://doi.org/10.1002/JCTB.5010020901>.
- 242 (11) Dette, H. P.; Qi, M.; Schröder, D. C.; Godt, A.; Koop, T. Glass-Forming
243 Properties of 3-Methylbutane-1,2,3-Tricarboxylic Acid and Its Mixtures with Water
244 and Pinonic Acid. *J. Phys. Chem. A* **2014**, *118* (34), 7024–7033.
245 <https://doi.org/10.1021/JP505910W>.
- 246 (12) Koop, T.; Bookhold, J.; Shiraiwa, M.; Pöschl, U. Glass Transition and Phase
247 State of Organic Compounds: Dependency on Molecular Properties and
248 Implications for Secondary Organic Aerosols in the Atmosphere. *Phys. Chem.*
249 *Chem. Phys.* **2011**, *13* (43), 19238–19255. <https://doi.org/10.1039/C1CP22617G>.
- 250 (13) Zobrist, B.; Marcolli, C.; Pedernera, D. A.; Koop, T. Do Atmospheric Aerosols
251 Form Glasses? *Atmospheric Chem. Phys.* **2008**, *8* (17), 5221–5244.
252 <https://doi.org/10.5194/ACP-8-5221-2008>.
- 253 (14) Asa-Awuku, A.; Engelhart, G. J.; Lee, B. H.; Pandis, S. N.; Nenes, A. Relating
254 CCN Activity, Volatility, and Droplet Growth Kinetics of β -Caryophyllene Secondary
255 Organic Aerosol. *Atmospheric Chem. Phys.* **2009**, *9* (3), 795–812.
256 <https://doi.org/10.5194/acp-9-795-2009>.
- 257 (15) Frosch, M.; Bilde, M.; Nenes, A.; Praplan, A. P.; Jurányi, Z.; Dommen, J.; Gysel,
258 M.; Weingartner, E.; Baltensperger, U. CCN Activity and Volatility of β -
259 Caryophyllene Secondary Organic Aerosol. *Atmospheric Chem. Phys.* **2013**, *13* (4),
260 2283–2297. <https://doi.org/10.5194/acp-13-2283-2013>.
- 261 (16) Smith, N. R.; Crescenzo, G. V.; Huang, Y.; Hettiyadura, A. P. S.; Siemens, K.; Li,
262 Y.; Faiola, C. L.; Laskin, A.; Shiraiwa, M.; Bertram, A. K.; Nizkorodov, S. A. Viscosity
263 and Liquid–Liquid Phase Separation in Healthy and Stressed Plant SOA. *Environ.*
264 *Sci. Atmospheres* **2021**, *1* (3), 140–153. <https://doi.org/10.1039/D0EA00020E>.
- 265 (17) Wang, J.; Shilling, J. E.; Liu, J.; Zelenyuk, A.; Bell, D. M.; Petters, M. D.;
266 Thalman, R.; Mei, F.; Zaveri, R. A.; Zheng, G. Cloud Droplet Activation of Secondary
267 Organic Aerosol Is Mainly Controlled by Molecular Weight, Not Water Solubility.
268 *Atmospheric Chem. Phys.* **2019**, *19* (2), 941–954. [https://doi.org/10.5194/acp-19-](https://doi.org/10.5194/acp-19-941-2019)
269 941-2019.

- 270 (18) Varutbangkul, V.; Brechtel, F. J.; Bahreini, R.; Ng, N. L.; Keywood, M. D.; Kroll, J.
271 H.; Flagan, R. C.; Seinfeld, J. H.; Lee, A.; Goldstein, A. H. Hygroscopicity of
272 Secondary Organic Aerosols Formed by Oxidation of Cycloalkenes, Monoterpenes,
273 Sesquiterpenes, and Related Compounds. *Atmospheric Chem. Phys.* **2006**, *6* (9),
274 2367–2388. <https://doi.org/10.5194/acp-6-2367-2006>.
- 275 (19) Zhao, D. F.; Buchholz, A.; Tillmann, R.; Kleist, E.; Wu, C.; Rubach, F.; Kiendler-
276 Scharr, A.; Rudich, Y.; Wildt, J.; Mentel, T. F. Environmental Conditions Regulate
277 the Impact of Plants on Cloud Formation. *Nat. Commun.* **2017**, *8* (1), 14067.
278 <https://doi.org/10.1038/ncomms14067>.
- 279 (20) Li, Y.; Pöschl, U.; Shiraiwa, M. Molecular Corridors and Parameterizations of
280 Volatility in the Chemical Evolution of Organic Aerosols. *Atmospheric Chem. Phys.*
281 **2016**, *16* (5), 3327–3344. <https://doi.org/10.5194/ACP-16-3327-2016>.
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